

EXPLANATION

UNCONSOLIDATED STRATA
QUATERNARY

INTRUSIVE ROCKS
MEDIOZIC

WISAHICKON GROUP

CAMBRO-ORDOVICIAN(?)
GLENHORN SUPERGROUP

PRECAMBRIAN

Diabase
Accumulations of rounded, rusty-weathering boulders and cobbles of a hard, massive, fine- to medium-grained, dark grey rock. Pyroxene and epidote are locally important, but dark green amphibole is locally important. Minor veins of amphibole and quartz also occur. A Mesozoic age is inferred due to the unmetamorphosed and unfoliated character of the rock.

Loch Raven Schist
L Loch Raven Schist. Uniform, medium to coarse-grained schist consisting of biotite-oligoclase-muscovite-quartz plus any combination of the following: chlorite, garnet, staurolite, kyanite, and tourmaline. Minor concordant and discordant quartz pods and stringers. Local concordant zones from decimeters to meters thick of a relatively quartz-feldspathic, micropoor schist. Thickness of Loch Raven Schist may exceed 2000 meters on the northeastern flank of the Phoenix Dome.

Cockeysville Marble
C Cockeysville Marble undivided. Low, hummocky-weathering outcrops and float; strongly silicified rock containing minor calcite and rarely pyroxene. Locally a yellow-brown-weathering, "honeycombed" appearing rock consisting of fine-grained silica pseudomorphing tremolite. Thickness ranges from 0.5 - 150 meters.

Setters Formation
S1 Garnet schist member. Medium-grained, garnet-biotite-muscovite-oligoclase-quartz schist, usually with accessory staurolite and/or kyanite, locally with tourmaline. Mesoscopically identical to the Loch Raven Schist. Thickness ranges from 0.5 - 225 meters.

S2 Gneiss member. Fine- to medium-grained muscovite-biotite-quartz-feldspathic gneiss with accessory garnet and tourmaline. Millimeter-scale laminations continuous over several meters. Local interlayers of biotite-rich schist and gneiss. Thickness ranges from 0.5 - 130 meters.

S3 Quartzite member. Fine- to medium-grained, centimeter-scale layered, flaggy weathering, muscovite-microcline quartzite, locally rich in tourmaline lying parallel to the bedding (?) planes. Locally, coarse quartzite and biotite schist. Thickness ranges from 0.5 - 550 meters, but the upper value may represent tectonic thickening in the northeast closure of the Phoenix Dome.

Q A single, roughly concordant lens of quite pure massive quartzite southeast of Phoenix.

S4 Conglomerate lens. Two lenses of microcline-poor, muscovite quartzite with abundant stretched quartz pebbles. Occurs at the base of the Setters Formation.

Baltimore Gneiss
DCB Layered gneiss member. Various layered biotite-quartz-feldspar gneiss containing muscovite, garnet, tourmaline, and rare sillimanite. Layering ranges from millimeter-scale and quite continuous to a more massive, quartz-feldspathic rock with discrete clots of biotite. Concordant to discordant pegmatites and quartz veins are common. Maximum thickness appears to be about 1500 meters, but due to complex folding, this figure is a rough estimate.

overprint Areas in which amphibole, interlayered with the above lithologies, comprises 20% of the gneiss. Fine- to medium-grained amphibole-biotite-plagioclase ± microcline and quartz gneiss.

* Stratigraphic nomenclature after Crowley, W. P., 1976. The geology of the crystalline rocks near Baltimore, and its bearing on the evolution of the Eastern Maryland Piedmont. Maryland Geol. Survey Rept. of Inv. no. 27, 40 p.

contact

approximated by distribution of rock exposures, float, and contrasts in soil composition and topography

anticline overturned anticline syncline overturned syncline

trace of the axial surface of F₂ folds.

trace of the axial surface of F₁ nappe structure (see schematic below), queried in areas where location is poorly known.

vertical inclined horizontal

S₁ foliation or schistosity (virtually parallel to compositional layering except in the noses of F₁ minor folds).

S₂ schistosity or cleavage (inclined) S₃ schistosity or fracture cleavage (inclined)

L₁ mineral elongation (rare), axis of minor F₁ fold, or intersection of compositional layering (bedding ?) with S₁

L₂ Mineral elongation, axis of minor F₂ fold of indeterminate symmetry, or intersection of compositional layering or S₁ with S₂ schistosity or cleavage

L₃ Axis of minor F₃ fold or intersection of compositional layering and/or S₁ with S₃ schistosity or fracture cleavage.

** horizontal linear structures are indicated with pointers on each end.

metamorphic zone boundary, approximately located, marks the first appearance of the mineral on the hatched side of the line in rocks of pelitic composition. Does not necessarily represent a specific metamorphic reaction.

Zone of retrograde metamorphism (indicated by fine-grained chlorite partially or wholly pseudomorphing garnet porphyroblasts).

Schematic three-dimensional representation of macroscopic folding in the Phoenix Quadrangle. Surface outlined by the solid line represents the contact between the Baltimore Gneiss and the Paleozoic sequence.

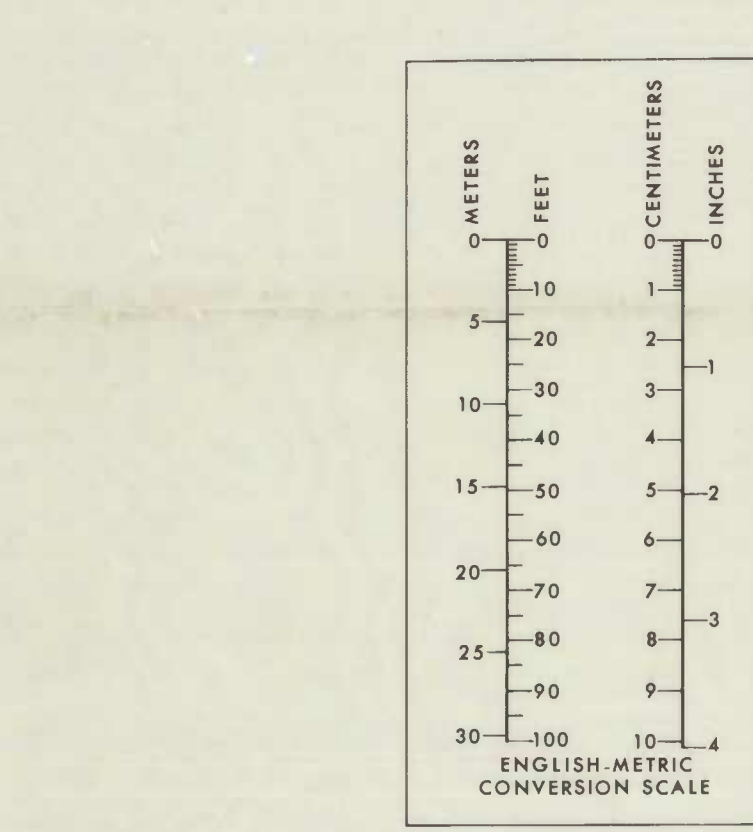
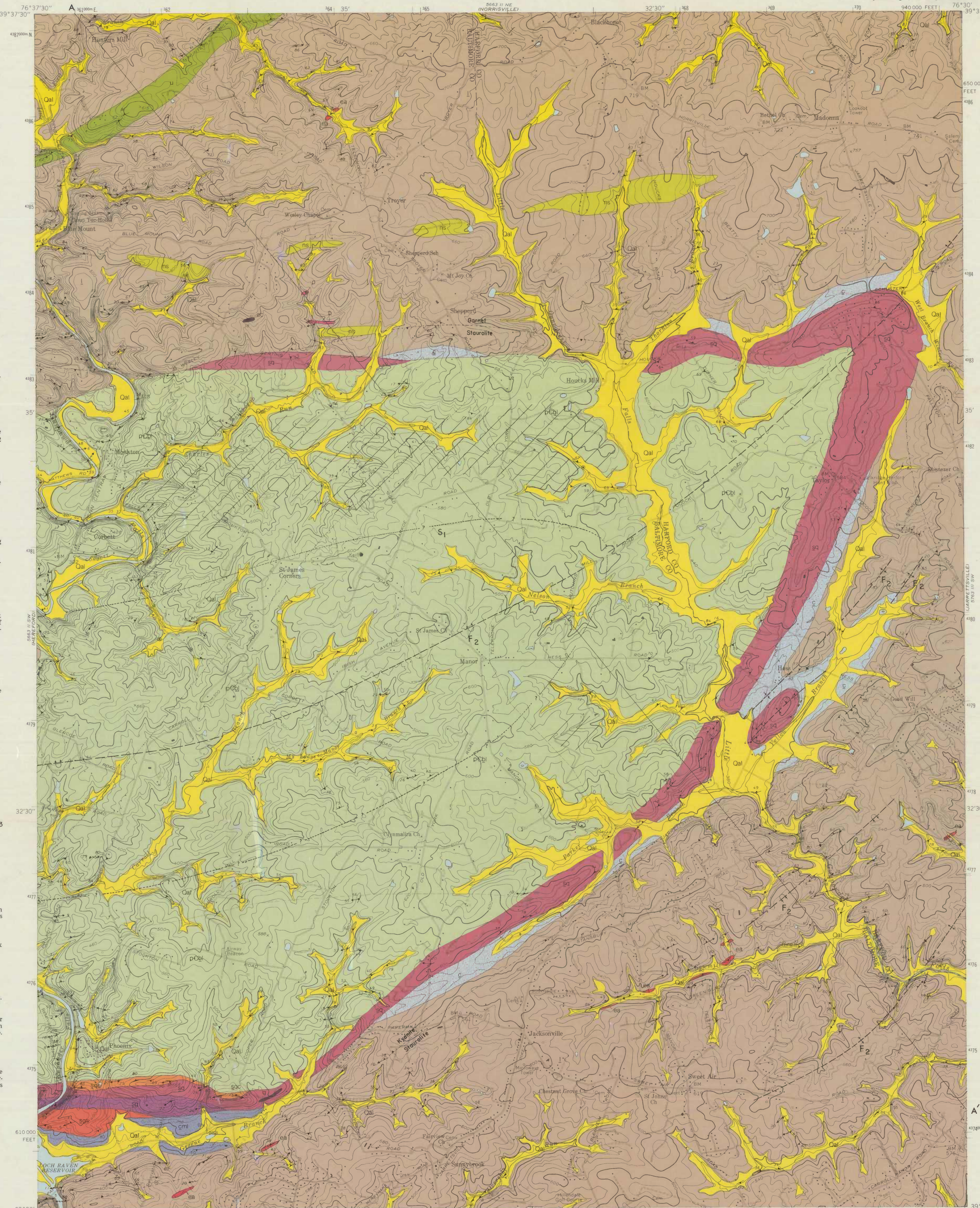
----- Axial surface of early (F₁) nappe (the Texas-Phoenix nappe of Crowley, 1976-8).

----- Representative axial surface of F₂ folding.

----- Representative axial surface of late (F₃) folding.

*** Crowley, W. P., 1976. The geology of the crystalline rocks near Baltimore and its bearing on the evolution of the Eastern Maryland Piedmont. Maryland Geol. Survey Rept. Inv. no. 27, 40 p.

† Editorial Note: Ultramafic rocks in the Loch Raven Schist on the north side of the Phoenix Dome (designated ns and u) are interpreted by Crowley*** (p. 26-27, Plate 1 and 2) to be allochthonous masses with thrust faults along their southern boundaries.

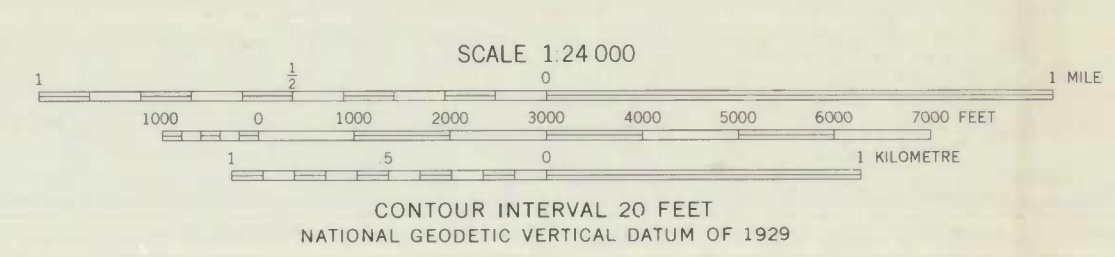


GEOLOGIC MAP OF THE PHOENIX QUADRANGLE, MARYLAND

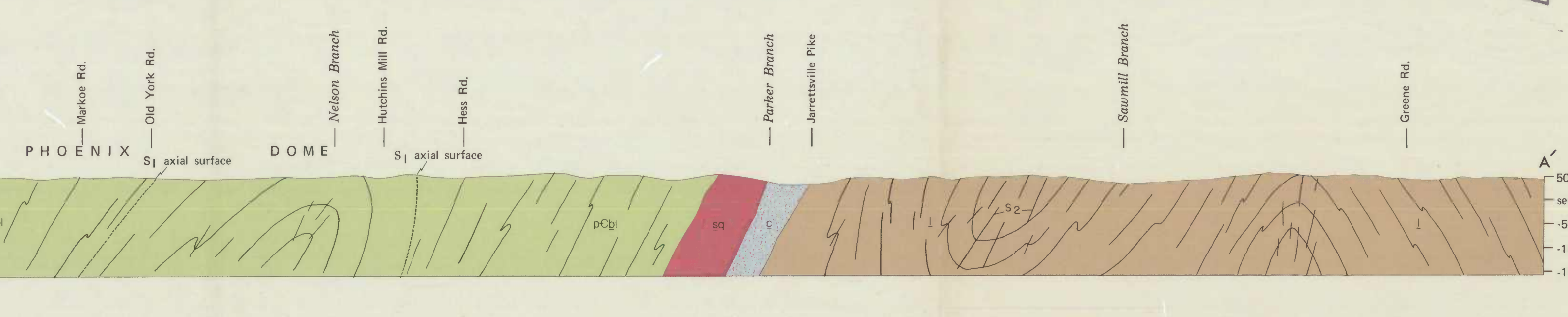
By
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1979

STATE OF MARYLAND
DEPARTMENT OF NATURAL RESOURCES
MARYLAND GEOLOGICAL SURVEY
Kenneth N. Weaver, Director

Copies of Map available from
Maryland Geological Survey
Johns Hopkins University
Baltimore, Maryland 21218



CROSS SECTION A-A'
Horizontal scale same as map scale, no vertical exaggeration.
Folded surface is S₁; F₂ Folds and S₂ Cleavage are Schematic



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